Complexity in Climate Change Manipulation Experiments

JUERGEN KREYLING AND CLAUS BEIER

Climate change goes beyond gradual changes in mean conditions. It involves increased variability in climatic drivers and increased frequency and intensity of extreme events. Climate manipulation experiments are one major tool to explore the ecological impacts of climate change. Until now, precipitation experiments have dealt with temporal variability or extreme events, such as drought, resulting in a multitude of approaches and scenarios with limited comparability among studies. Temperature manipulations have mainly been focused only on warming, resulting in better comparability among studies. Congruent results of meta-analyses based on warming experiments, however, do not reflect a better general understanding of temperature effects, because the potential effects of more complex changes in temperature, including extreme events, are not yet covered well. Heat, frost, seasonality, and spatial variability in temperature are ecologically important. Embracing complexity in future climate change experiments in general is therefore crucial.

Keywords: climate change, experiments, warming, temporal variability, extreme events

climate change will alter the fundamental conditions of future terrestrial ecosystems, with temperature and precipitation being the most affected (Solomon et al. 2007). Climate manipulation experiments are one important tool for understanding the response of ecosystems to such changes (e.g., Rustad 2008). Temperature and precipitation have different characteristics, which lead to different considerations regarding the design of experiments and the scenarios to be tested.

Precipitation is temporally and spatially variable, and future scenarios are relatively uncertain (Solomon et al. 2007). Recently, Beier and colleagues (2012) summed up the current state of the art in precipitation manipulation experiments and pointed out several crucial aspects to be considered in the future—for example, biased geographic coverage, artifacts related to rain-out shelter design, and the need for proper controls. One of their basic insights was that the available precipitation manipulations have been carried out in many different contexts, through many different designs, and they have been used to test very different (and often simplistic) precipitation scenarios. Therefore, these experiments are hardly comparable, which has led to a lack of formal meta-analyses. Beier and colleagues (2012) argued that this might be because changes in precipitation regimes are more complex and uncertain than those in temperature, which makes their scenarios more difficult to define and the required range of experimental conditions more complex and less comparable. Focusing on more simple and uniform scenarios of change has been suggested as a strategy that would lead to better comparability (Knapp et al. 2012, Fraser et al. 2013) and has been the basis for meta-analyses (Wu et al. 2011).

Temperature is a continuous variable, and the climate-driven changes in temperature are less variable and more predictable than are those of precipitation. In experimentation, this apparent homogeneity in temperature relative to precipitation has led to the application of approaches almost entirely focused on average increases in temperature (figure 1). In consequence, the conducted warming experiments are more comparable, which facilitates comprehensive meta-analyses (e.g., Rustad et al. 2001, Lin et al. 2010, Dieleman et al. 2012). However, the focus on gradual and positive shifts of the mean temperature (figure 1; Jentsch et al. 2007) is accompanied by a lack of studies on extremes and temporal variability in these experiments. Consequently, complexity is not yet adequately reflected in temperature change experiments.

Locally and regionally, the global temperature increase will manifest through highly varied changes, including uneven warming and, in some places, even cooling and shortor long-term extreme temperature changes such as heat waves, increased freeze—thaw cycles, and winter warm spells (Solomon et al. 2007). Such changes may have large impacts on ecosystem processes and function. For instance, single heat waves can alter plant community compositions (White et al. 2000). Likewise, minimum temperature events during

BioScience 63: 763–767. ISSN 0006-3568, electronic ISSN 1525-3244. © 2013 by American Institute of Biological Sciences. All rights reserved. Request permission to photocopy or reproduce article content at the University of California Press's Rights and Permissions Web site at www.ucpressjournals.com/reprintinfo.asp. doi:10.1525/bio.2013.63.9.12

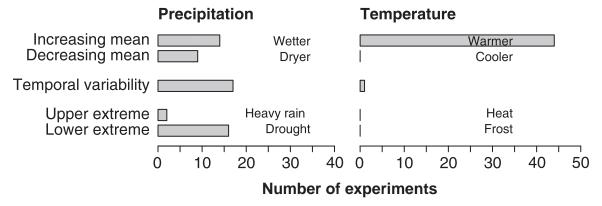


Figure 1. Complexity is embraced in precipitation experiments, whereas temperature experiments so far show a clear focus on rising mean temperatures (warming). Temporal variability includes increased or reduced variability in temperature or precipitation over time. The data include 45 temperature manipulation experiments and 43 precipitation experiments taken from the TERACC (Terrestrial Ecosystem Response to Atmospheric Climatic Change; www.umaine.edu/teracc) and INTERFACE (Integrated Network for Terrestrial Ecosystem Research on Feedbacks to the Atmosphere and ClimatE; www.bio.purdue.edu/INTERFACE) databases. Manipulations of more than one category were applied in several experiments; individual experiments could therefore be counted more than once.

winter (Kreyling et al. 2012a) or spring (Gu et al. 2008) can create lethal stress, can alter the competitive balance within plant communities, and can affect biogeochemical cycles (Mulholland et al. 2009). The ecological and evolutionary importance of minimum temperature events in the context of climate change was summed up by Inouye (2000).

Until now, treatments were applied in most warming experiments only during the growing season (Rustad et al. 2001), whereas climate projections clearly indicate differences in warming trends between seasons, with strongest shifts expected for winter (Christensen et al. 2007). Warming during different seasons causes contrasting effects, with winter and spring warming more important than summer warming for plant phenology and productivity in cold, northern ecosystems (Aerts et al. 2006). An earlier onset of the growing season in response to global warming can furthermore increase the risk of frost damage (Inouye 2008, Augspurger 2013). Temperature variability over days or a few weeks is also ecologically relevant but not well investigated. Warming pulses of a few days over winter can lead to massive dieback in tundra vegetation (Bokhorst et al. 2009), and the effects of warming pulses may benefit some and be a detriment to other plant functional types in temperate zones (Kreyling et al. 2010). Changes in the frequency of freeze-thaw cycles can furthermore increase carbon loss from ecosystems and can affect microbial communities (Larsen et al. 2002).

The scientific community has applied different experimental approaches for precipitation and temperature manipulations: Precipitation experiments have been focused mostly on single events (in some cases, extreme events) and complexity, whereas temperature manipulations have been focused on shifts in mean conditions (figure 1). This is understandable, because precipitation is a discrete and

stochastic variable, whereas warming is a continuous variable. The availability of meta-analyses and their congruent results for temperature change (Rustad et al. 2001, Lin et al. 2010, Dieleman et al. 2012), however, should not be mistaken as a sign of a general understanding of temperature effects on terrestrial ecosystems, because these experiments cover only average warming and not the potential effects associated with extreme events and more complex changes in temperature. The examples presented above imply that temperature change also contains ecologically important challenges in regard to complexity (e.g., extremes, temporal variability).

For precipitation changes, Knapp and colleagues (2008) demonstrated that increased variability and extremity are likely to become crucial factors controlling ecological effects in terrestrial ecosystems at all moisture levels, especially because these changes will lead to increased frequency of threshold exceedance. Beier and colleagues (2012) further showed how these complexities need to be systematically addressed in future precipitation experiments. Similar arguments can be raised for temperature, and the examples given above clearly demonstrate that the complexity of temperature shifts is ecologically important. For temperature, there is therefore a clear need to embrace complexity in future, well-designed experiments.

In comparison with the study of chronic changes in mean conditions, several challenges arise for experimental designs focused on the complexity of climatic drivers (table 1). Projections of magnitudes and frequencies of occurrence of extreme events are generally uncertain, which complicates the choice of a single "correct" scenario to be experimentally tested. Here, gradient or regression-type experiments (Beier et al. 2012) appear useful to determine response surfaces rather than single responses and to identify the thresholds of

Table 1. Advantages and challenges for climate change manipulation experiments.			
Factors	Focus	Advantages	Challenges
Single factor	Chronic shift of mean conditions	Sound projections from climate models available, few scenarios, easy to replicate across systems	To test generalization with multifactor and multisite experiments across biomes, to test gradients of different manipulation strengths
	Pulsed shifts in temporal variability and single (extreme) events	Extremes determine the exceedance of ecological thresholds and mortality	There are no sound projections from climate models available to test for thresholds (gradient or regression designs); sensitivity can be system specific, so generalization across sites might be elusive
Multifactor	Chronic shifts of mean conditions	More realistic than single-factor manipulations (multifactor experiments tend to level responses out among factors)	Selection of factor combinations (the number of combinations more than doubles for any new factor added)
	Chronic shifts of mean conditions combined with pulsed shifts in temporal variability and single (extreme) events	More realistic in relation to future scenarios	The unlimited number of scenarios (test system sensitivity to single factors first, use gradient or regression designs, couple with modeling, focus on process understanding)

sensitivities. These sensitivities, however, will differ among systems, just as limiting factors differ. Furthermore, the suite of possible scenarios for the different factors and their combinations are numerous, and identification of a "correct" scenario appears impossible. Therefore, such experiments should focus, rather, on process understanding first, before any step toward broader generalization is made. Seeking generalization by repeating manipulations across different systems (the multisite approach; Beier et al. 2004, Knapp et al. 2012, Fraser et al. 2013) is an important step forward in experimental climate impact assessments but is clearly limited by logistical constraints on the complexity of possible manipulations. We suggest that experiments on complexity should be initially focused on single factors at various strengths, intensities, or frequencies, should be carried out at single or a few sites, and should ideally be closely coupled with process-based modeling from the start in order to generate hypotheses, guide the choice of scenarios and the responses to be measured, and generalize the results beyond the site- and scenario-specific conditions. The link between virtual experiments with unlimited choices of scenarios using such models (e.g., Gerten et al. 2008, Luo et al. 2008) and the direct verification by process-based experiments appears to be crucial.

Some experimental approaches have allowed for the study of extreme rainfall events (e.g., drought, heavy rainfall) and variability of precipitation (Fay et al. 2000, Beier et al. 2004, Jentsch et al. 2007), but challenges arise with regard to the simulation of temperature extremes and temperature variability in field experiments because of trade-offs among the demand for more intense heating, technical possibilities, and artifacts. Passive warming systems lack sufficient control to obtain high temperature increases for long periods of time (Bruhn et al. 2012), whereas greenhouses or chambers entail significant unwanted side effects on temperature, light, and wind (e.g., Rasmussen et al. 2002). Infrared heating treatments can simulate warm spells and heat waves, but they require large amounts of energy and, therefore, financial

resources and involve obvious constraints for application in large areas and with tall vegetation (De Boeck and Nijs 2011, De Boeck et al. 2012). Cooling is even more challenging in the field. In ecosystems with predictable winter snow cover, snow removal can be used to create cold extremes (e.g., Kreyling et al. 2012a). During other seasons and in systems without snow cover, portable devices for the simulation of air frost have been suggested (Thorpe et al. 1993). Alternatively, "realistic" experiments in the field may be combined with laboratory and chamber studies of specific processes in plants and mesocosms (e.g., Kreyling et al. 2012b) or with long-term monitoring in which naturally occurring extremes are analyzed (see, e.g., Ciais et al. 2005 as an example of the ecological consequences of an extreme heat wave). These alternatives clearly have drawbacks, because laboratory and chamber studies are associated with the abovementioned artifacts and with a lack of ecosystem focus, and long-term monitoring data series will miss true controls or references and may not include the monitoring of relevant responses because of the unplanned nature of the extreme events.

Experiments on single extreme events and on long-term trends in mean conditions differ in the time scale needed for meaningful manipulations but not in that needed for meaningful quantification of the ecological responses. Grassland community shifts, for instance, take about 10 years to reach a new quasiequilibrium in response to alterations in the precipitation regime (Heisler and Weltzin 2006). Climate manipulations concerning mean conditions therefore need to be continued for several years to allow for sound investigations of their effects. Early responses may be transient and may lead to improper conclusions regarding long-term responses (Hollister et al. 2005). Similarly, experiments on single events should follow the effects over time to understand recovery, adaptation, and long-term consequences, which might differ from short-term effects (Kreyling et al. 2010). Although manipulations might be carried out just once over very short time periods, the effects of repeated

events over many years (e.g., Sowerby et al. 2008), as well as hysteresis and alternative stable states (Scheffer and Carpenter 2003) or ecological memory (Walter et al. 2013), are important aspects to be monitored in the long run.

Finally, climate factors work in combination. The few examples of combinations of average temperature, precipitation, and carbon dioxide (CO₂) change, however, led to divergent results that are not always predictable on the basis of the individual effects (Shaw et al. 2002, Larsen et al. 2011, Dieleman et al. 2012). Multifactor experiments are therefore crucial in order to test interactions among several simultaneous factors. In such experiments, the inclusion of temporal variability in drivers such as seasonality and extreme events will inevitably increase complexity considerably. In addition, nonclimatic drivers need to be taken into account, such as land-use change, biodiversity loss, soil structure change, nitrogen deposition or increased levels of atmospheric CO₂. Combining experiments and gradient studies by conducting the same experimental manipulations along environmental gradients can improve the external validity of controlled experiments (Arft et al. 1999, Peñuelas et al. 2007, Beier et al. 2008, 2012).

We advocate the incorporation of complexity into climate change experiments, especially with respect to temperature change. The almost exclusive focus on increases in mean temperature needs to be broadened to studies of seasonality, variability, and extreme events in order to allow for a sound understanding of the ecological implications of climate change. In addition to recently proposed multisite experiments across the globe focused on simple changes in mean climatic conditions (Fraser et al. 2013), we recommend testing the sensitivity of various systems to temperature and precipitation extremes and variability using gradient or regression approaches in single-factor experiments. Process-based modeling based on the results of these simple experiments will allow for virtual experiments on combined drivers in different settings for which the main findings need to be verified by more complex multifactor experiments.

Acknowledgments

We thank three anonymous reviewers for constructive critique on earlier versions of the manuscript. JK received financial support from the German Academic Exchange Service and CB from the Villum Kann Rasmussen Foundation's CLIMAITE project, the European Union's INCREASE network (contract no. 227628), and the Analysis and Experimentation on Ecosystems project (contract no. 312690).

References cited

- Aerts R, Cornelissen JHC, Dorrepaal E. 2006. Plant performance in a warmer world: General responses of plants from cold, northern biomes and the importance of winter and spring events. Plant Ecology 182: 65–77.
- Arft AM, et al. 1999. Responses of tundra plants to experimental warming: Meta-analysis of the International Tundra Experiment. Ecological Monographs 69: 491–511.

- Augspurger CK. 2013. Reconstructing patterns of temperature, phenology, and frost damage over 124 years: Spring damage risk is increasing. Ecology 94: 41–50.
- Beier C, et al. 2004. Novel approaches to study climate change effects on terrestrial ecosystems in the field: Drought and passive nighttime warming. Ecosystems 7: 583–597.
- . 2008. Carbon and nitrogen cycles in European ecosystems respond differently to global warming. Science of the Total Environment 407: 692–697.
- . 2012. Precipitation manipulation experiments—Challenges and recommendations for the future. Ecology Letters 15: 899–911.
- Bokhorst SF, Bjerke JW, Tømmervik H, Callaghan TV, Phoenix GK. 2009. Winter warming events damage sub-Arctic vegetation: Consistent evidence from an experimental manipulation and a natural event. Journal of Ecology 97: 1408–1415.
- Bruhn D, Larsen KS, de Dato GD, Duce P, Zara P, Beier C, Schmidt IK, Clausen S, Mikkelsen N. 2012. Improving the performance of infrared reflective night curtains for warming field plots. Agricultural and Forest Meteorology 173: 53–62.
- Christensen JH, et al. 2007. Regional climate projections. Pages 847–940 in Solomon S, Qin D, Manning M, Marquis M, Averyt K, Tignor MMB, Miller HL Jr, Chen Z, eds. Climate Change 2007: The Physical Science Basis. Cambridge University Press.
- Ciais P, et al. 2005. Europe-wide reduction in primary productivity caused by the heat and drought in 2003. Nature 437: 529–533.
- De Boeck HJ, Nijs I. 2011. An alternative approach for infrared heater control in warming and extreme event experiments in terrestrial ecosystems. Journal of Ecology 99: 724–728.
- De Boeck HJ, Kimball BA, Miglietta F, Nijs I. 2012. Quantification of excess water loss in plant canopies warmed with infrared heating. Global Change Biology 18: 2860–2868.
- Dieleman WIJ, et al. 2012. Simple additive effects are rare: A quantitative review of plant biomass and soil process responses to combined manipulations of ${\rm CO_2}$ and temperature. Global Change Biology 18: 2681–2693.
- Fay PA, Carlisle JD, Knapp AK, Blair JM, Collins SL. 2000. Altering rainfall timing and quantity in a mesic grassland ecosystem: Design and performance of rainfall manipulation shelters. Ecosystems 3: 308–319.
- Fraser LH, et al. 2013. Coordinated distributed experiments: An emerging tool for testing global hypotheses in ecology and environmental science. Frontiers in Ecology and the Environment 11: 147–155.
- Gerten D, et al. 2008. Modelled effects of precipitation on ecosystem carbon and water dynamics in different climatic zones. Global Change Biology 14: 2356–2379.
- Gu L, Hanson PJ, Post WM, Kaiser DP, Yang B, Nemani R, Pallardy SG, Meyers T. 2008. The 2007 Eastern US spring freeze: Increased cold damage in a warming world? BioScience 58: 253–262.
- Heisler JL, Weltzin JF. 2006. Variability matters: Towards a perspective on the influence of precipitation on terrestrial ecosystems. New Phytologist 172: 189–192.
- Hollister RD, Webber PJ, Bay C. 2005. Plant response to temperature in Northern Alaska: Implications for predicting vegetation change. Ecology 86: 1562–1570.
- Inouye DW. 2000. The ecological and evolutionary significance of frost in the context of climate change. Ecology Letters 3: 457–463.
- 2008. Effects of climate change on phenology, frost damage, and floral abundance of montane wildflowers. Ecology 89: 353–362.
- Jentsch A, Kreyling J, Beierkuhnlein C. 2007. A new generation of climate change experiments: Events, not trends. Frontiers in Ecology and the Environment 5: 365–374.
- Knapp AK, et al. 2008. Consequences of more extreme precipitation regimes for terrestrial ecosystems. BioScience 58: 811–821.
- 2012. Past, present, and future roles of long-term experiments in the LTER network. BioScience 62: 377–389.
- Kreyling J, Beierkuhnlein C, Jentsch A. 2010. Effects of soil freeze-thaw cycles differ between experimental plant communities. Basic and Applied Ecology 11: 65–75.

- Kreyling J, Haei M, Laudon H. 2012a. Absence of snow cover reduces understory plant cover and alters plant community composition in boreal forests. Oecologia 168: 577–587.
- Kreyling J, Peršoh D, Werner S, Benzenberg M, Wöllecke J. 2012b. Short-term impacts of soil freeze-thaw cycles on roots and root-associated fungi of *Holcus lanatus* and *Calluna vulgaris*. Plant and Soil 353: 19–31.
- Larsen KS, Jonasson S, Michelsen A. 2002. Repeated freeze—thaw cycles and their effects on biological processes in two arctic ecosystem types. Applied Soil Ecology 21: 187–195.
- Larsen KS, et al. 2011. Reduced N cycling in response to elevated CO₂, warming, and drought in a Danish heathland: Synthesizing results of the CLIMAITE project after two years of treatments. Global Change Biology 17: 1884–1899.
- Lin D, Xia J, Wan S. 2010. Climate warming and biomass accumulation of terrestrial plants: A meta-analysis. New Phytologist 188: 187–198.
- Luo Y, et al. 2008. Modeled interactive effects of precipitation, temperature and [CO₂] on ecosystem carbon and water dynamics in different climatic zones. Global Change Biology 14: 1986–1999.
- Mulholland PJ, Roberts BJ, Hill WR, Smith JG. 2009. Stream ecosystem responses to the 2007 spring freeze in the southeastern United States: Unexpected effects of climate change. Global Change Biology 15: 1767–1776.
- Peñuelas J, et al. 2007. Response of plant species richness and primary productivity in shrublands along a north–south gradient in Europe to seven years of experimental warming and drought: Reductions in primary productivity in the heat and drought year of 2003. Global Change Biology 13: 2563–2581.
- Rasmussen L, Beier C, Bergstedt A. 2002. Experimental manipulations of old pine forest ecosystems to predict the potential tree growth effects of increased CO₂ and temperature in a future climate. Forest Ecology and Management 158: 179–188.
- Rustad LE. 2008. The response of terrestrial ecosystems to global climate change: Towards an integrated approach. Science of the Total Environment 404: 222–235.
- Rustad LE, Campbell JL, Marion GM, Norby RJ, Mitchell MJ, Hartley AE, Cornelissen JHC, Gurevitch J, GCTE-NEWS. 2001. A metaanalysis of the response of soil respiration, net nitrogen mineralization,

- and aboveground plant growth to experimental ecosystem warming. Oecologia 126: 543–562.
- Scheffer M, Carpenter SR. 2003. Catastrophic regime shifts in ecosystems: Linking theory to observation. Trends in Ecology and Evolution 18: 648–656.
- Shaw MR, Zavaleta ES, Chiariello NR, Cleland EE, Mooney HA, Field CB. 2002. Grassland responses to global environmental changes suppressed by elevated CO₂. Science 298: 1987–1990.
- Solomon S, Qin D, Manning M, Marquis M, Averyt K, Tignor MMB, Miller HL Jr, Chen Z, eds. 2007. Climate Change 2007: The Physical Science Basis. Cambridge University Press.
- Sowerby A, Emmett BA, Tietema A, Beier C. 2008. Contrasting effects of repeated summer drought on soil carbon efflux in hydric and mesic heathland soils. Global Change Biology 14: 2388–2404.
- Thorpe PC, MacGillivray CW, Priestman GH. 1993. A portable device for the simulation of air frosts at remote field locations. Functional Ecology 7: 503–505.
- Walter J, Jentsch A, Beierkuhnlein C, Kreyling J. 2013. Ecological stress memory and cross stress tolerance in plants in the face of climate extremes. Environmental and Experimental Botany. (18 June 2013; www.sciencedirect.com/science/article/pii/S0098847212000482) doi: 10.1016/j.envexpbot.2012.02.009
- White TA, Campbell BD, Kemp PD, Hunt CL. 2000. Sensitivity of three grassland communities to simulated extreme temperature and rainfall events. Global Change Biology 6: 671–684.
- Wu Z, Dijkstra P, Koch GW, Peñuelas J, Hungate BA. 2011. Responses of terrestrial ecosystems to temperature and precipitation change: A meta-analysis of experimental manipulation. Global Change Biology 17: 927–942.

Juergen Kreyling (juergen.kreyling@uni-bayreuth.de) is affiliated with the Bayreuth Center of Ecology and Environmental Research (BayCEER), Department of Biogeography, at the University of Bayreuth, in Bayreuth, Germany. Claus Beier is affiliated with the Department of Chemical and Biochemical Engineering at the Technical University of Denmark, in Roskilde, and the Norwegian Institute for Water Research (NIVA) in Oslo.